



Preliminary Results of recent Decadal Predictions with EC-Earth

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Recent really means recent

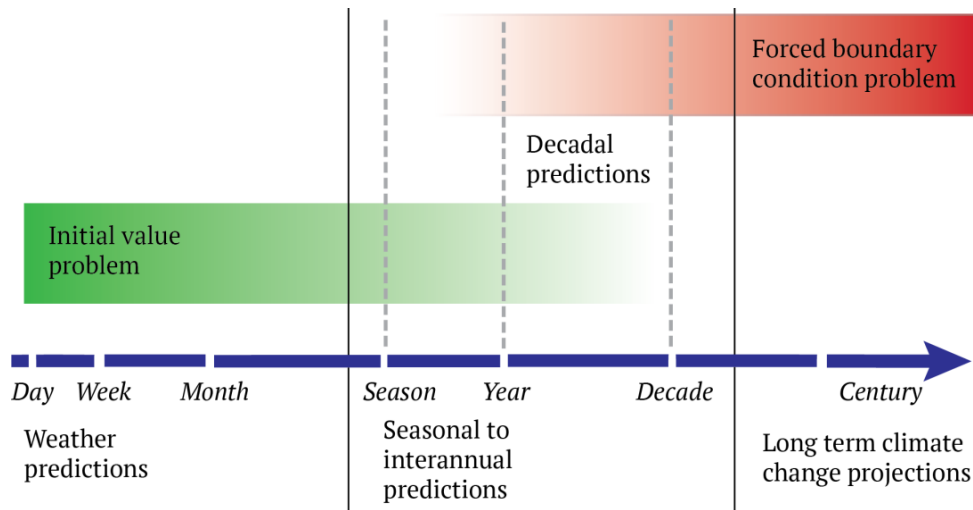
1. Decadal Climate Prediction Project – DCPD
2. Model Setup + Initialisation
3. First Quality Assessment of Decadal Hindcast Set
4. Climate Sensitivity in Historical Experiments
5. What am I actually doing here?

Near Term Climate Prediction



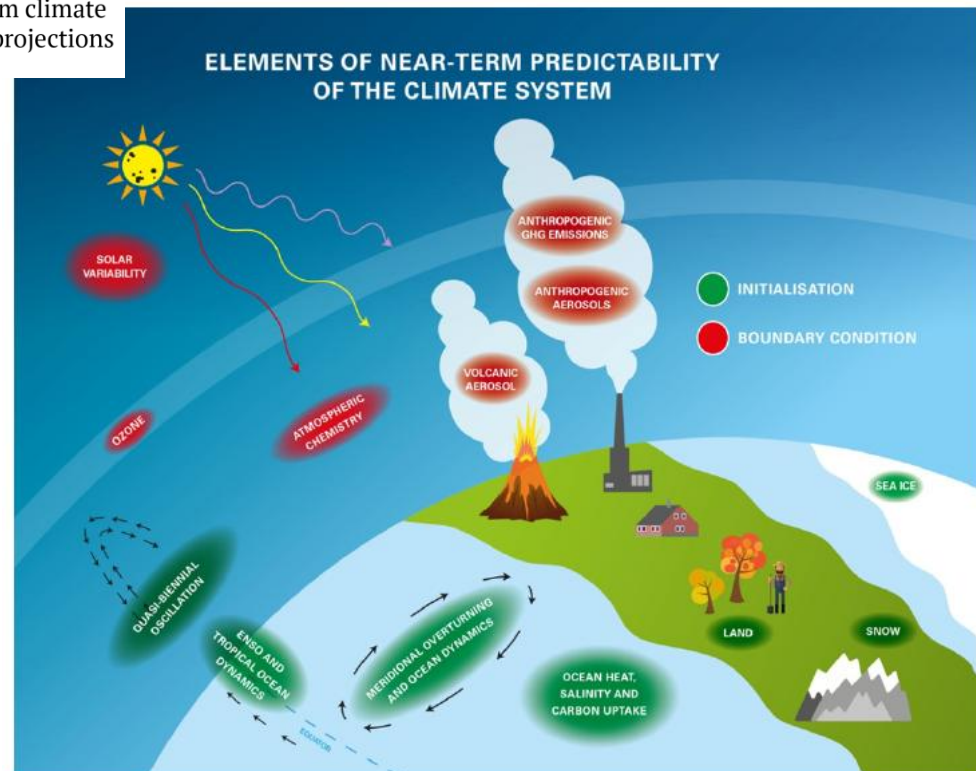
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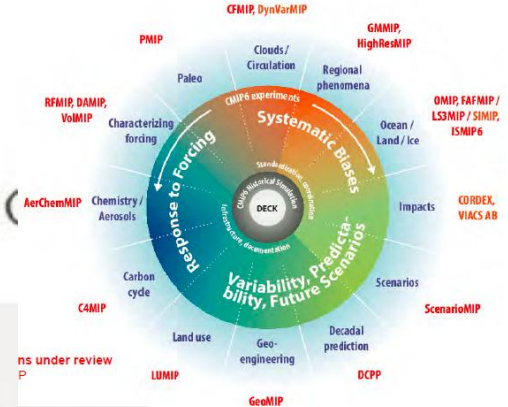
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Kushnir et al. 2018

Meehl et al. 2009





The Decadal Climate Prediction Project (DCPP)

The term "decadal prediction" encompasses predictions on annual, multi-annual to decadal timescales. The possibility of making skilful forecasts on these timescales and the ability to do so is investigated by means of predictability studies and retrospective predictions (hindcasts) made using the current generation of climate models and by empirical methods. Skilful decadal prediction of relevant climate parameters is a Key Deliverable of the WCRP's Grand Challenge

of Near-term Climate Prediction

The DCP envisions three components:

- A** ◦ **Hindcasts:** the design and organization of a coordinated decadal prediction (hindcast) component of CMIP6 in conjunction with the seasonal prediction and climate modelling communities
- B** ◦ **Forecasts:** the ongoing production of experimental quasi-operational decadal climate predictions in support of multi-model annual to decadal forecasting and the application of the forecasts
- C** ◦ **Predictability, mechanisms and case studies:** the organization and coordination of decadal climate predictability studies and of case studies of particular climate shifts and variations including the study of the mechanisms that determine these behaviours

DCPP A Protocol

- yearly initialised Experiments 1960-2017
- 10 members
- 10 years

Group/Model	Institution	Country	Component		
			A	B	C
EC-Earth	BSC/SMHI	Spain/Sweden	Y/M	Y/M	M/M
GFDL	NOAA	USA	Y	Y	M
FIO-ESM	FIO	China	Y	Y	M
NUIST-CSM	IPRC	USA/China	Y	Y	N
BCC	BCC	China	Y	N	N
CAS-ESM	CAS	China	M	M	M
MIROC	JAMSTEC/JMA	Japan	Y	y	Y
Can-ESM5	CCCma	Canada	Y	Y	M
CNRM-CERFACS	CNRS	France	M	M	Y
MetOffice	MetOffice	UK	Y	Y	Y
Ureading/Stat	UReading	UK	Y	Y	Y
IPSL	LOCEAN	France	Y	N	Y
NERCS/NorCPS	GRI	Norway	Y	Y	Y
CMCC	CMCC	Italy	M	M	Y
MPG	MPI	Germany	Y	Y	Y
NCAR/CESM	NCAR	USA	Y	Y	Y
BESM	INPE	Brazil	Y	M	M
FGOALS	IAP	China	Y	M	Y
INM	RAS	Russia	Y	N	Y

20 Groups

- 13 Countries
- Component A (hindcasts)
 - 16 Yes, 4 Maybe
- Component B (forecasts)
 - 11 Yes, 6 Maybe, 3 No
- Component C (mechanisms)
 - 11 Yes, 7 Maybe, 2 No

EC-Earth 3.3

yearly start dates

1960-2017, starting 1st Nov

10 members

Model:

Atmosphere: IFS, T255L91

Ocean: NEMO, ORCA1L75

Sea ice: LIM 3

+ OASIS coupler

Initial conditions, full-field

Atmosphere:

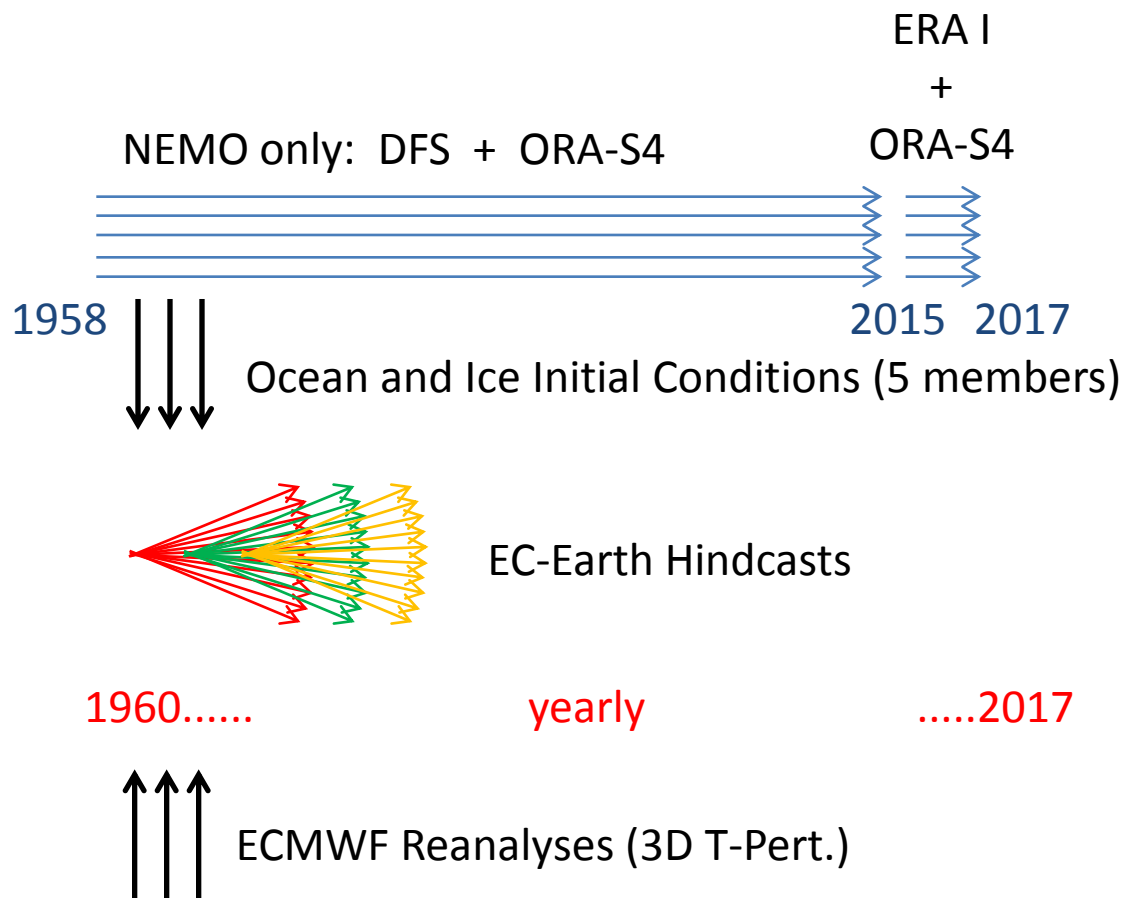
ECMWF reanalyses (ERAInt + ERA40)

with GPCP-corrected land surface

Ocean and sea-ice reconstruction:

ICs are produced using a NEMO only simulation

forced by DFS (ERA-I) fluxes and nudged towards ORA S4



Following Guemas et al. (Clim. Dyn., 2014)

ATM:

Interpolated to
model grid with
OpenIFS
(now performed
locally at BSC)

Atmosphere
reanalysis
(ERA 40 + Interim)

Land reanalysis
(ERA-Land)

Ocean reanalysis
(ORAS4)

Sea Ice
reanalysis

Produced at BSC

LAND:

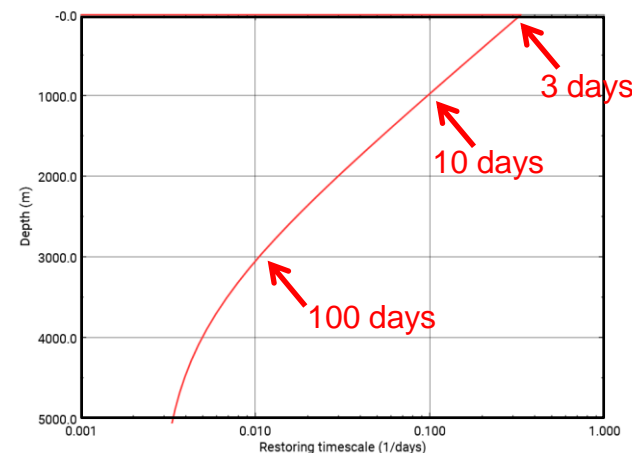
Offline land-surface
simulation with
near-surface
meteorology and
corrected fluxes
from ERA-Interim

OCE+ SI:

Historical reconstruction with NEMO-LIM stand
alone, forced with ERA-40/Interim fluxes, and
nudged globally towards 3D T and S from ORAS4

$$\left[\begin{array}{l} \text{Default surface} \\ \text{restoring coefficients} \\ \gamma_T = -40 \text{ W/m}^2/\text{K} \\ \gamma_S = -150 \text{ kg/m}^2/\text{s/psu} \end{array} \right]$$

Default 3D restoring timescales

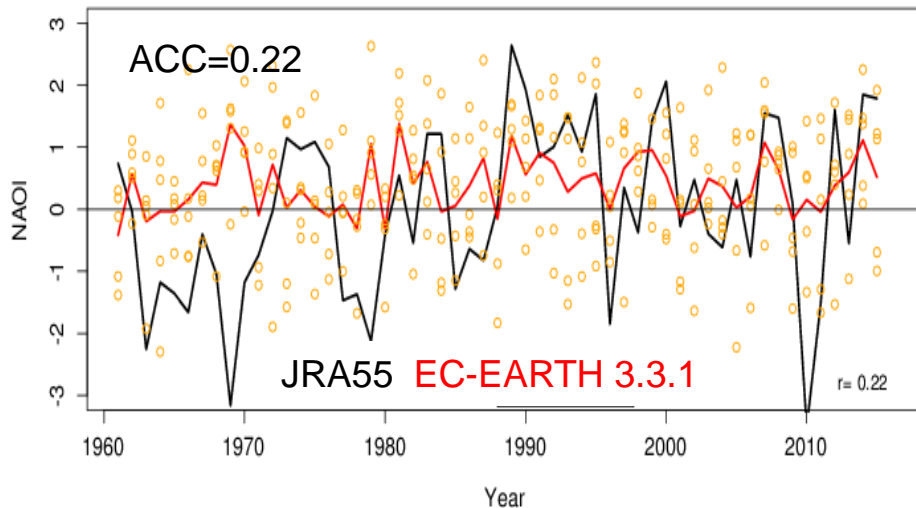


Predictability of atmospheric variables

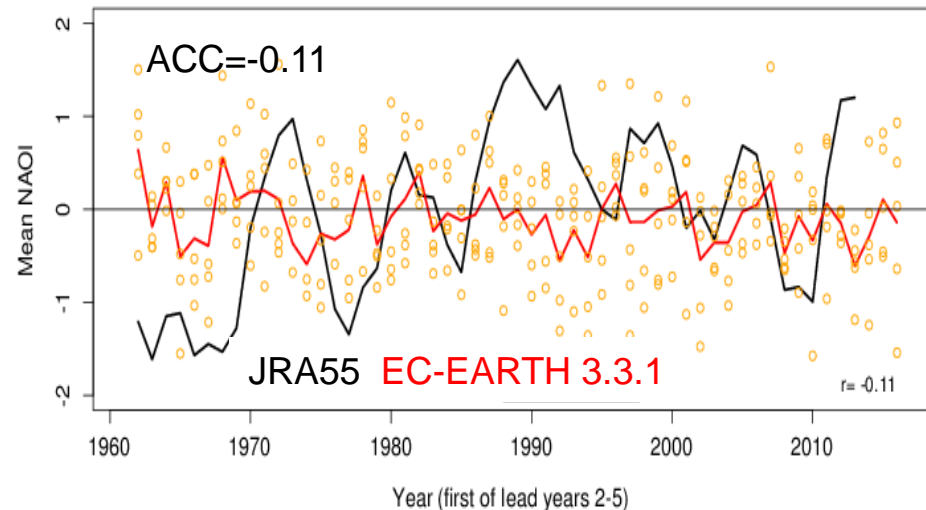
NAO-Index as defined in Jianping & Wang (2003):

35N-65N difference of standardised zonally averaged winter MSLP anomalies in the longitudinal band 80W-40E

NAO DJF Year 1 (forecast months 2-4)



NAO DJF Years 2-5



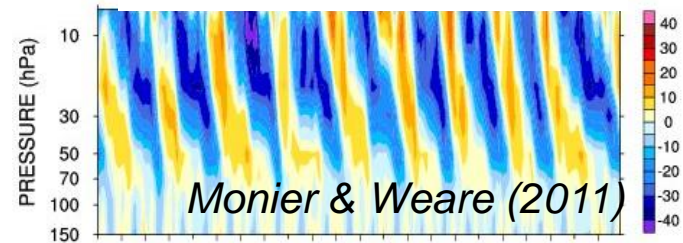
Unlike in other prediction systems, no significant skill in the NAO is obtained
Larger ensemble sizes might be needed to detect the predictable NAO component

Predictability of atmospheric variables

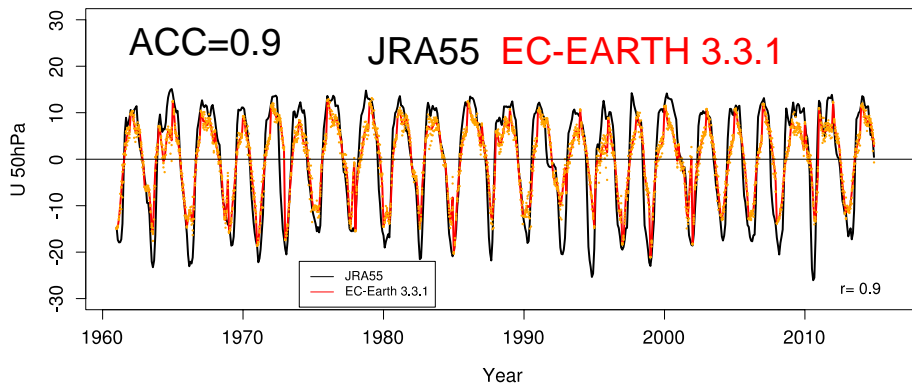
QBO-Index defined as the zonally averaged zonal wind at the **Equator and 50hPa**

Anomalies are computed with respect to the seasonal mean (as a function of lead time)

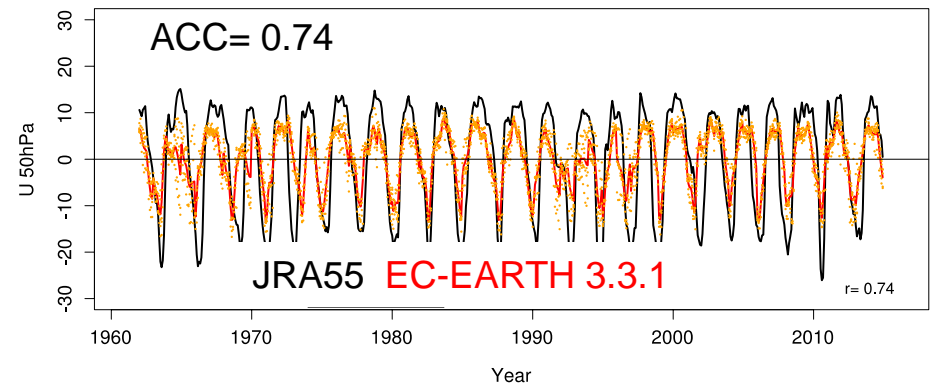
Equatorial Zonal Wind (m/s)



QBO in forecast year 1 (months 1-12)



QBO in forecast year 2



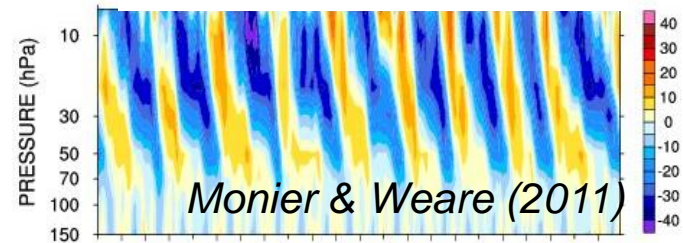
The QBO seems to be highly predictable in the first and second forecast years
This skill is unrelated to the effect of the seasonal cycle

Predictability of atmospheric variables

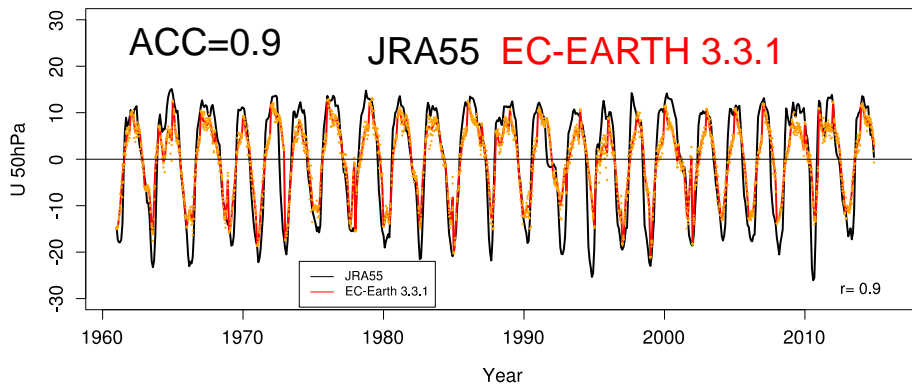
QBO-Index defined as the zonally averaged zonal wind at the **Equator and 50hPa**

Anomalies are computed with respect to the seasonal mean (as a function of lead time)

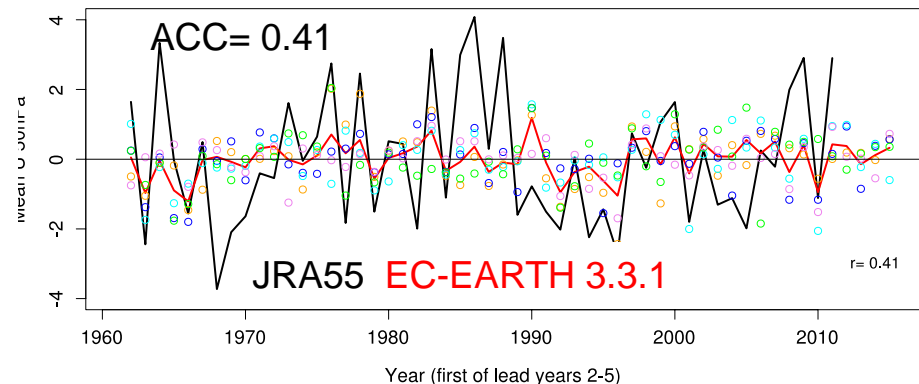
Equatorial Zonal Wind (m/s)



QBO in forecast year 1 (months 1-12)



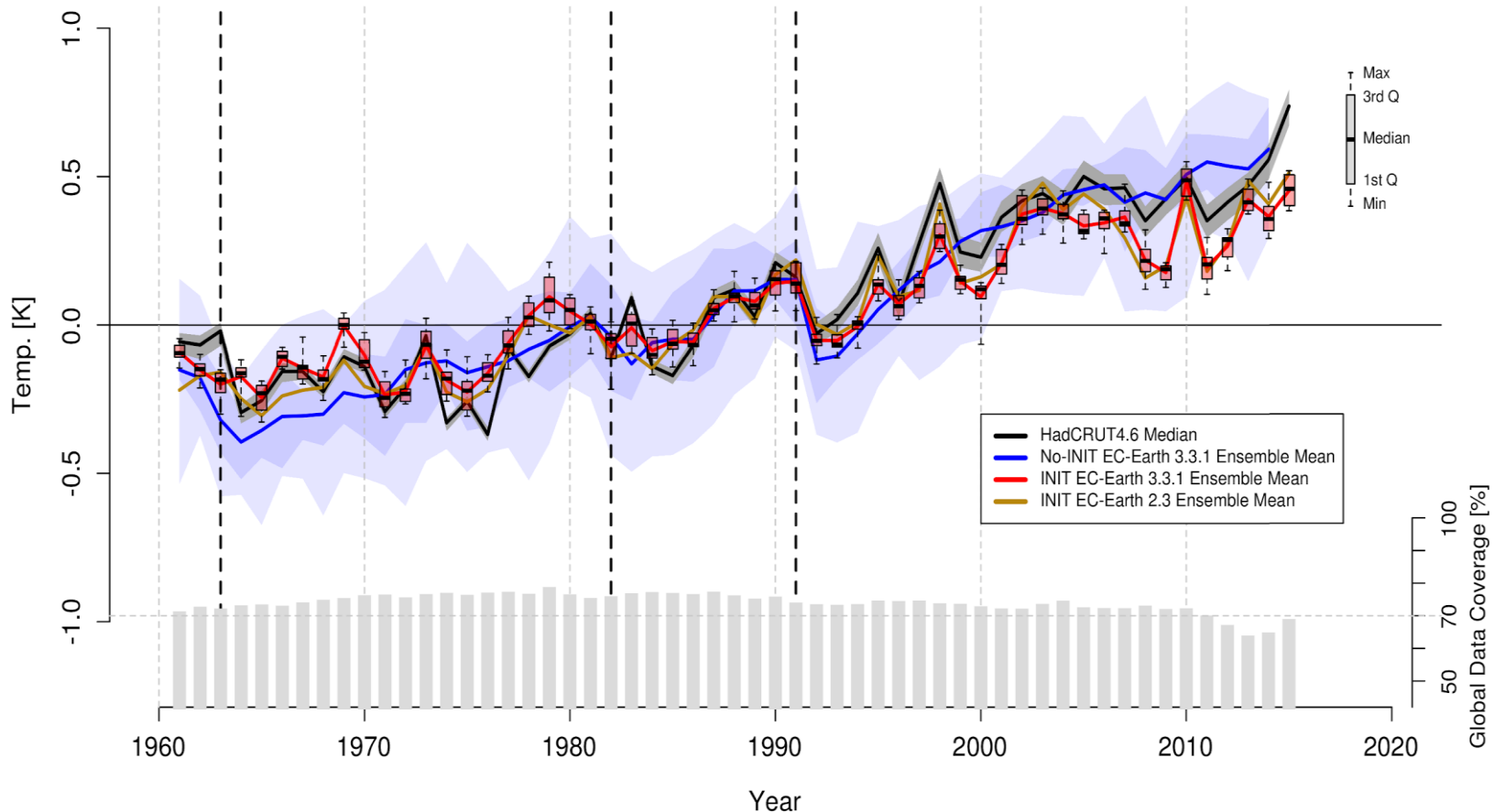
QBO in forecast years 2-5



The QBO seems to be partly predictable from the second to fifth forecast years
Skill is unrelated to the seasonal cycle, and is a consequence of its strong periodicity

Forecast Year 1 (M3-14)

Combination of 2m temperature over land and SST over ocean

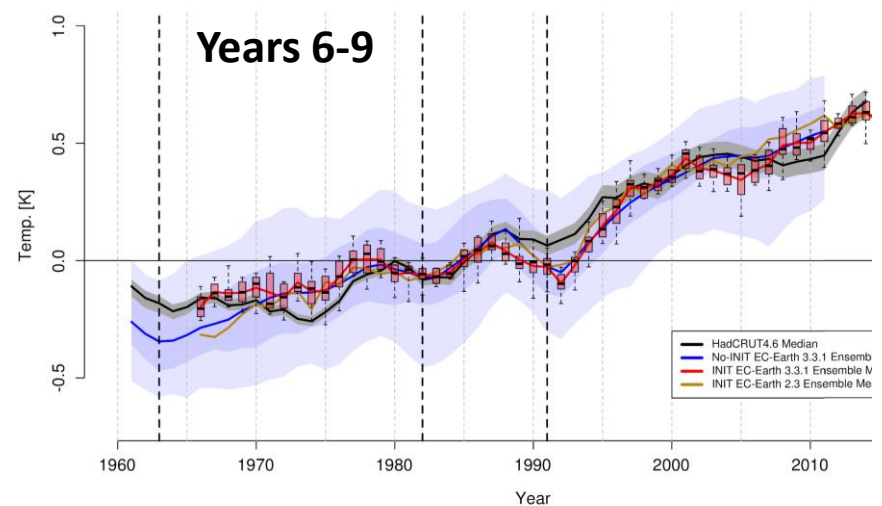
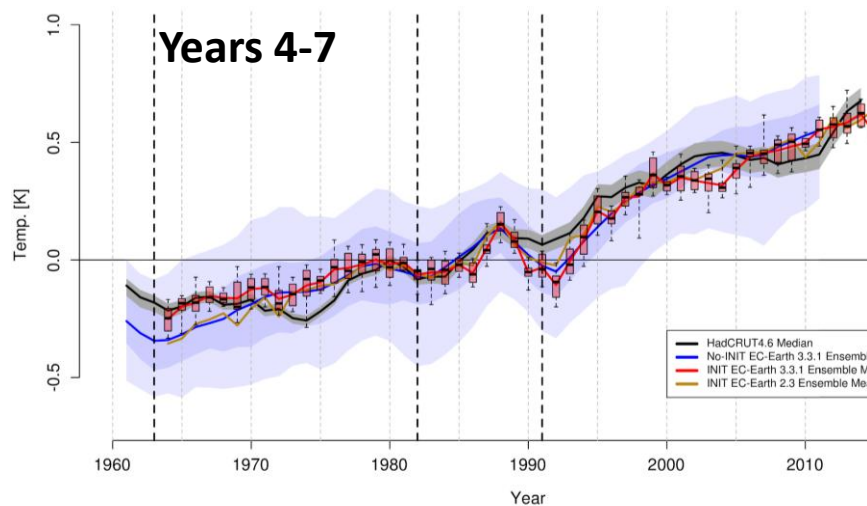
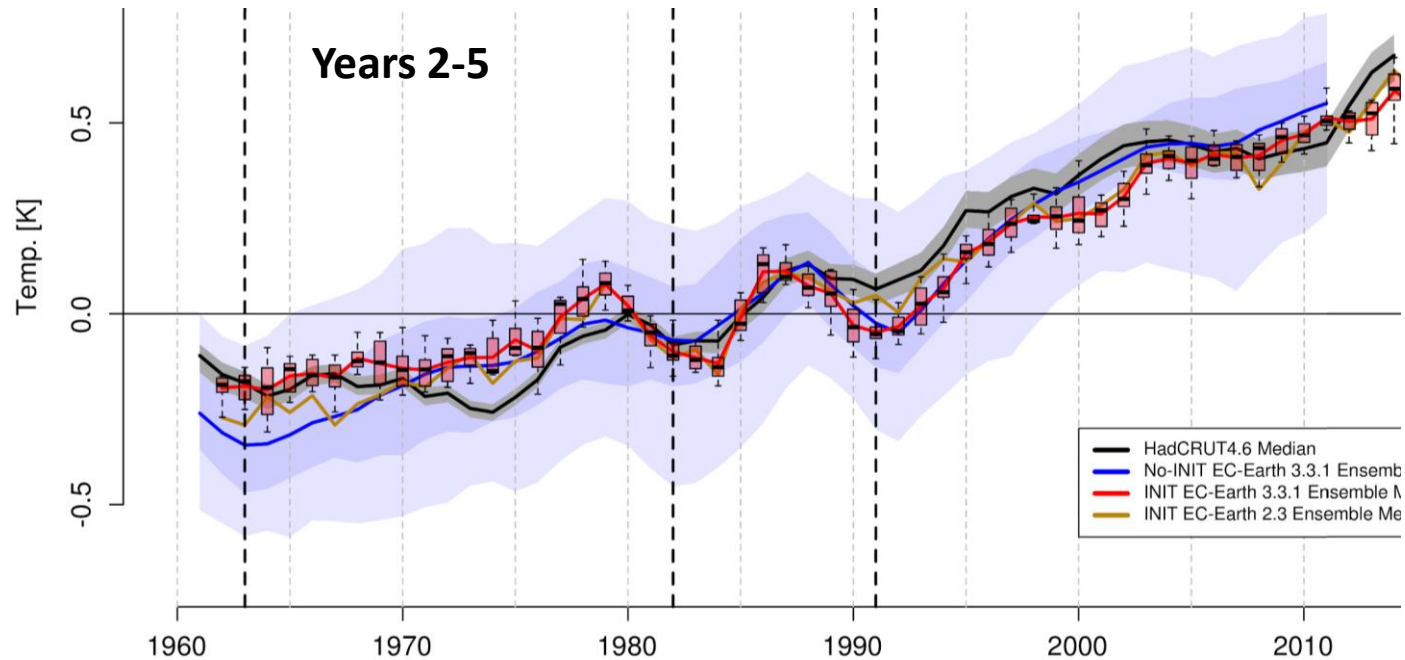


Global Mean Surface Temperature



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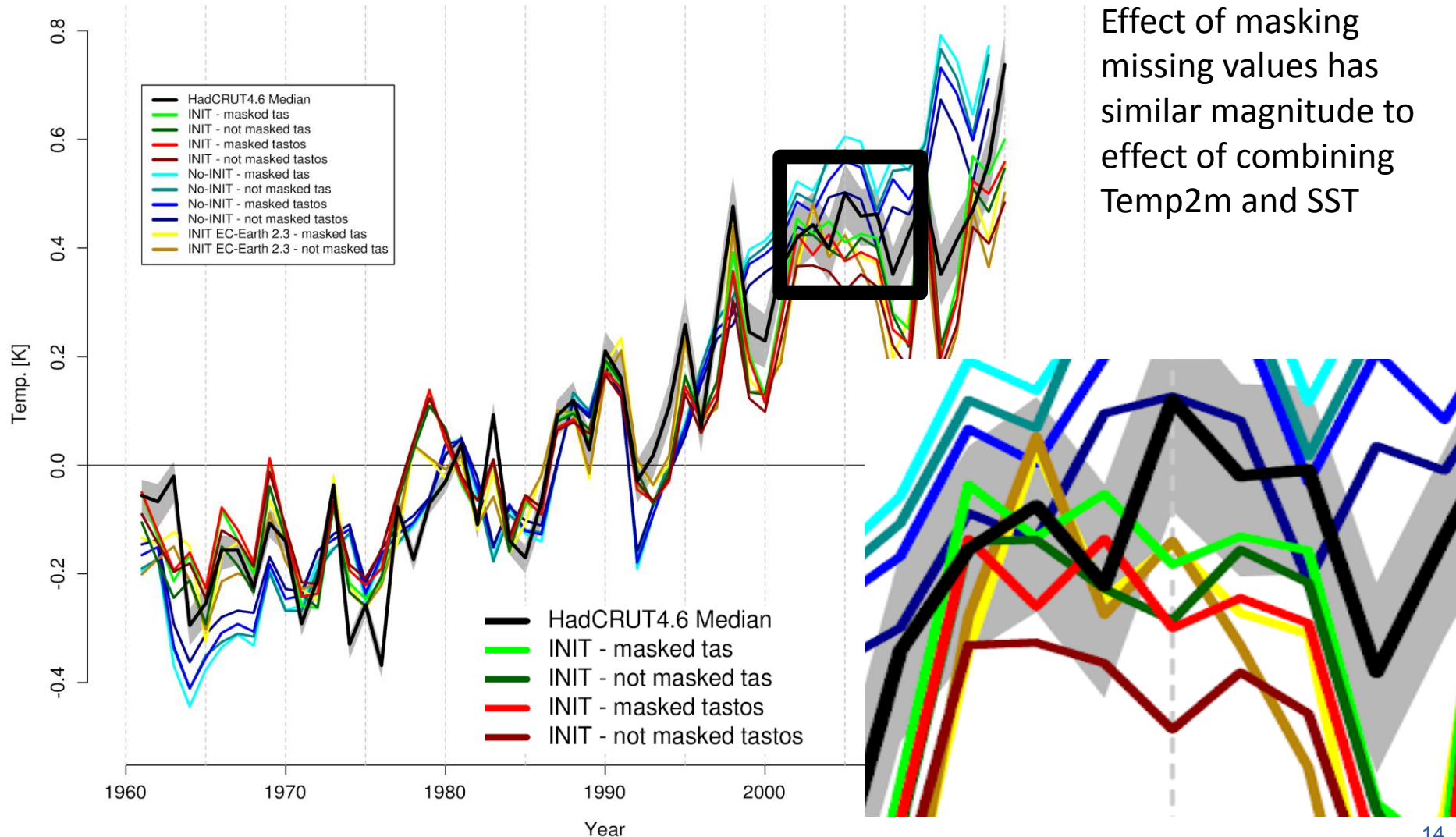
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Effect of combining 2m Temp and SST



Global Mean Surface Temperature, Forecast Year 1

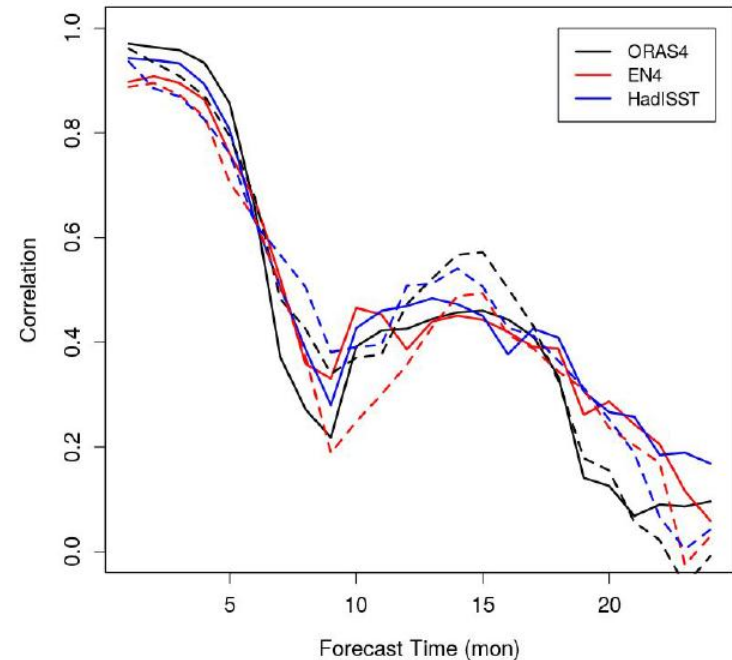
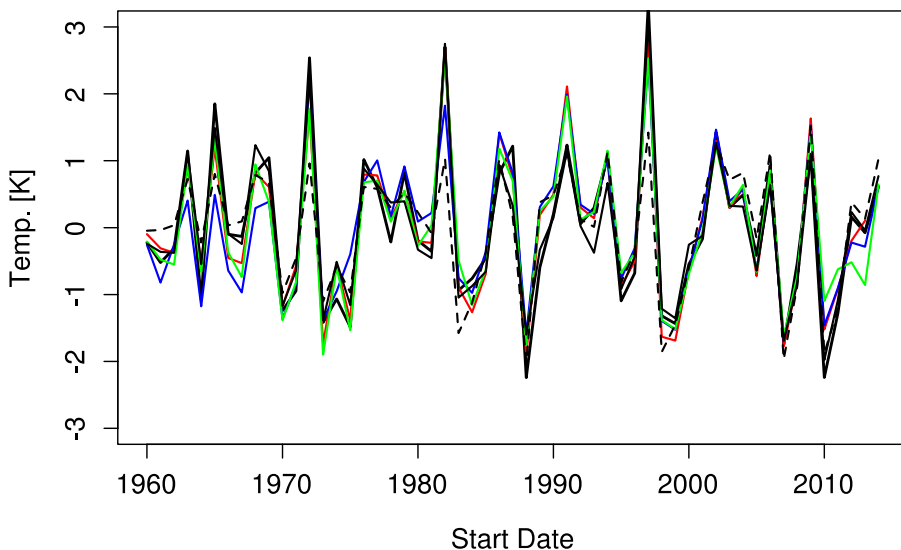


Predictability of ocean variables

ENSO-Index: Niño3.4

Regional average of SST anomalies in the box [5N-5S, 170W-120W]

Observed and Forecasted Niño3.4



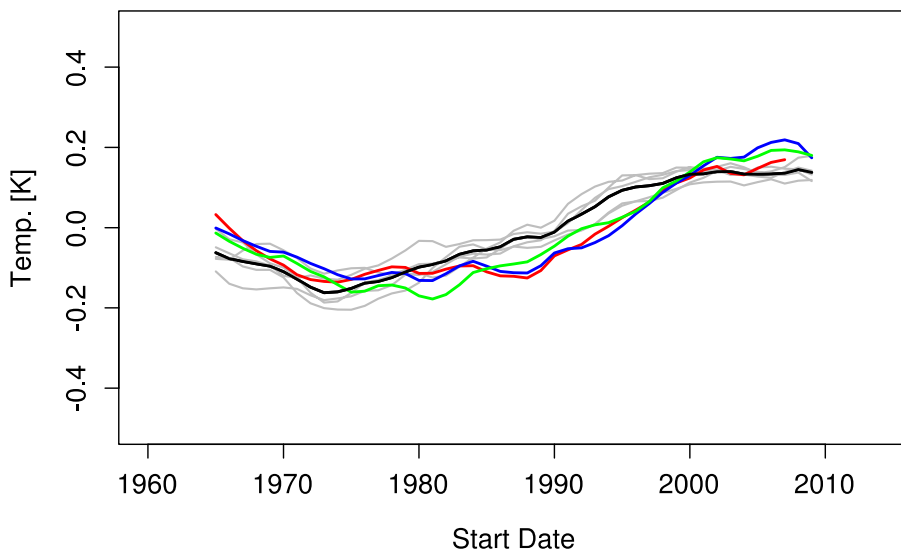
The hindcast maintains high ENSO skill during the first 5 forecast months (DEC through APR), and experiences some skill reemergence in the second winter

Predictability of ocean variables

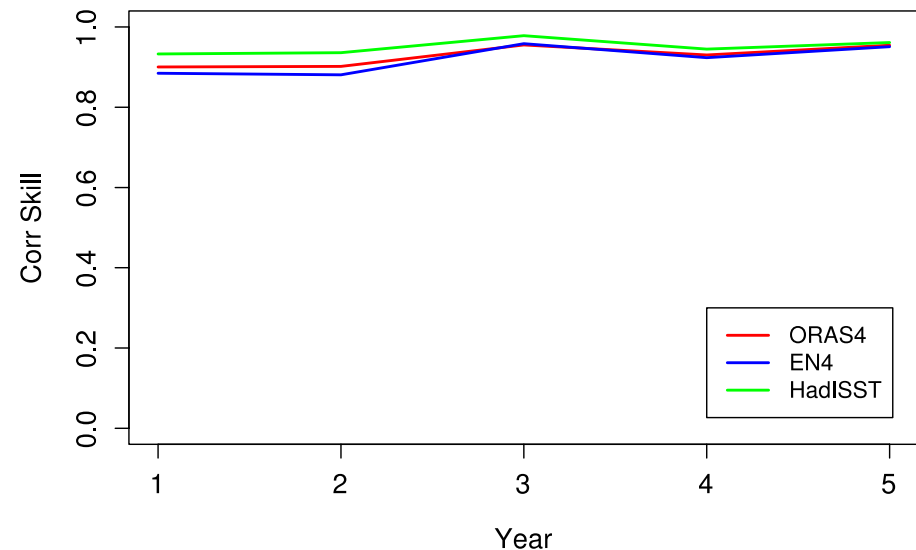
AMV-Index as defined in Trenberth & Shea (2006):

Decadally smoothed difference between the SST average in the **North Atlantic [0-60N, 80W-0]** and the **global mean SST**

Observed and Forecasted AMV



AMV Anomaly Correlation Coefficient



The models has really high AMV skill, at least for the first 5 forecast years

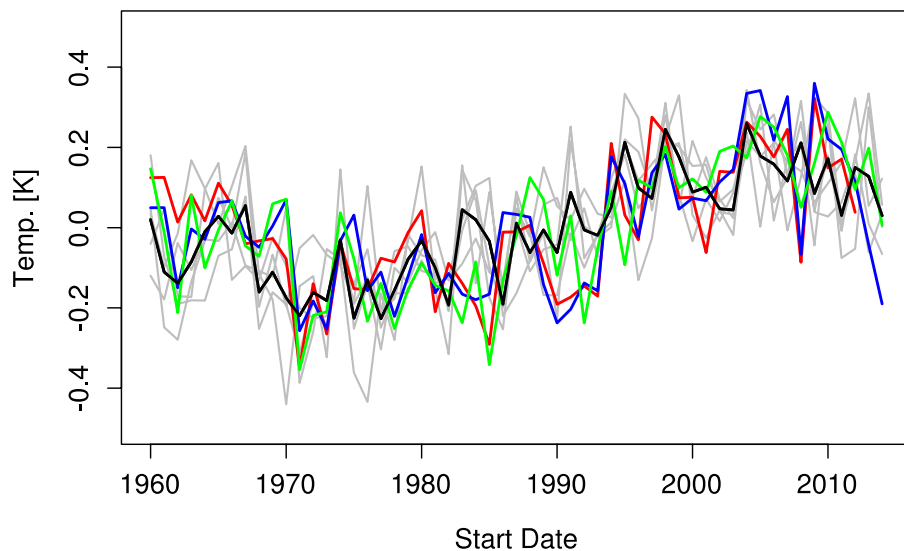
It might be due to the reduced degrees of freedom as a result of the smoothing

Predictability of ocean variables

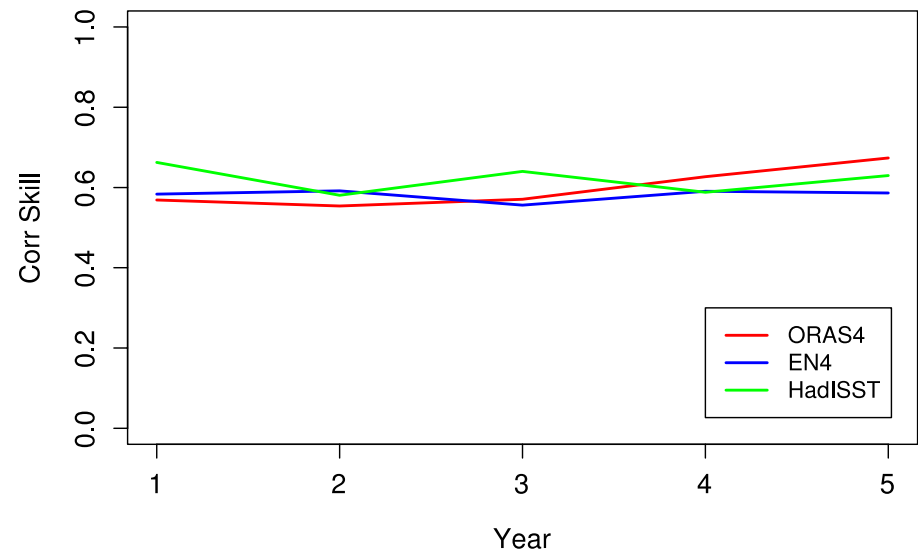
AMV-Index as defined in Trenberth & Shea (2006):

Raw difference between the SST average in the **North Atlantic [0-60N, 80W-0]** and the **global mean SST**

Observed and Forecasted AMV



AMV Anomaly Correlation Coefficient

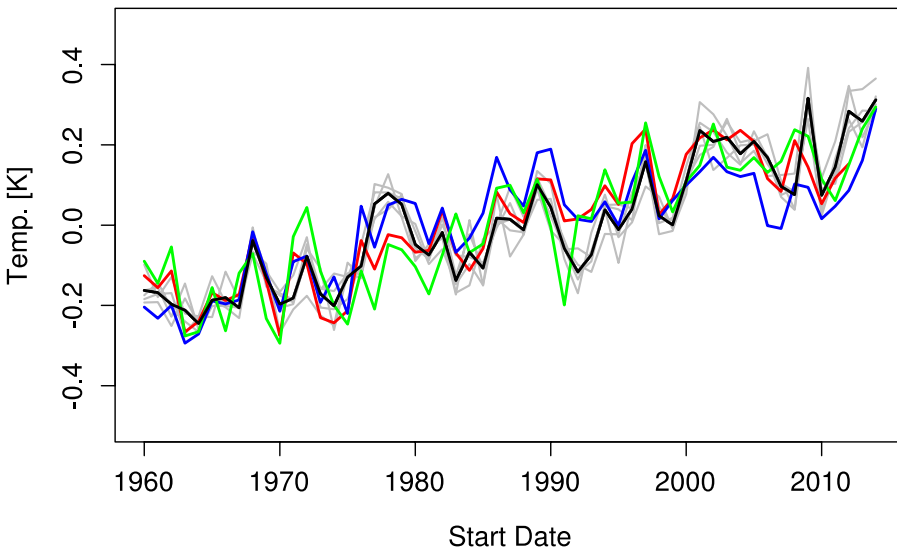


The skill drops but remains high also for the unfiltered timeseries

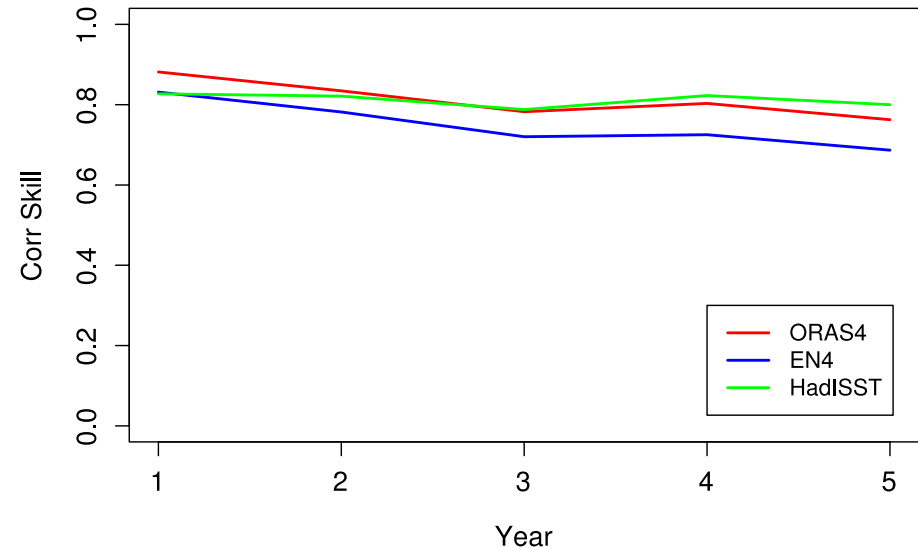
Predictability of ocean variables

Global Mean SST

Global Mean SST (1st Year)



ACC Global Mean SST

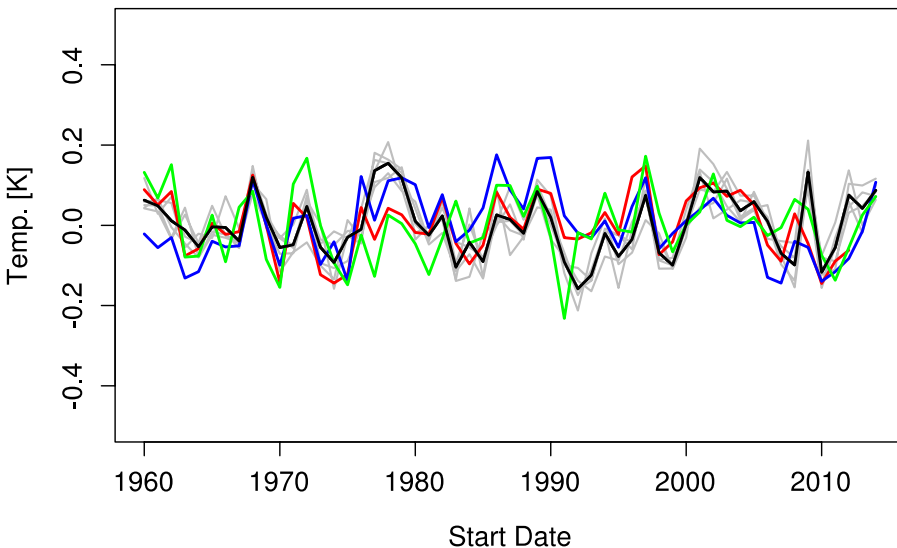


Global mean temperature is also largely predictable, although a non-negligible contribution comes from the long-term trend.

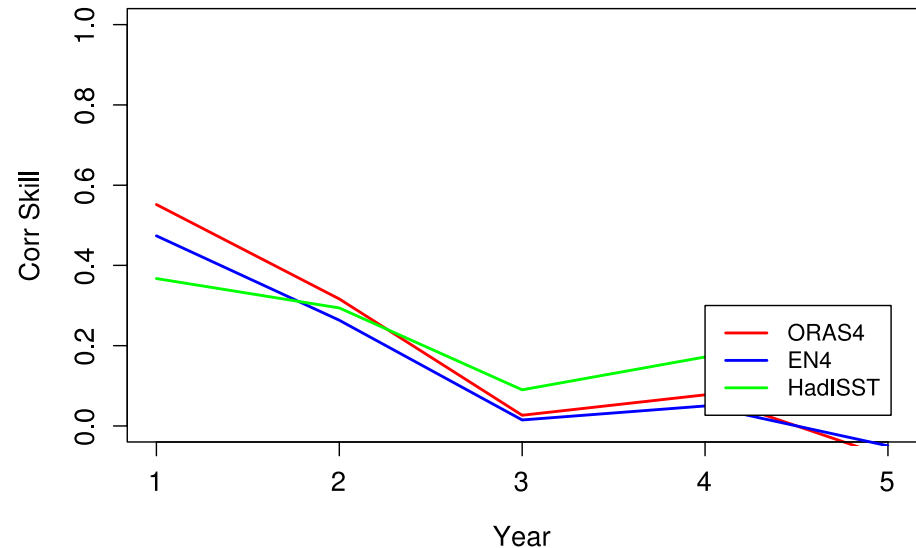
Predictability of ocean variables

Global Mean SST

Global Mean SST (1st Year) Detrended

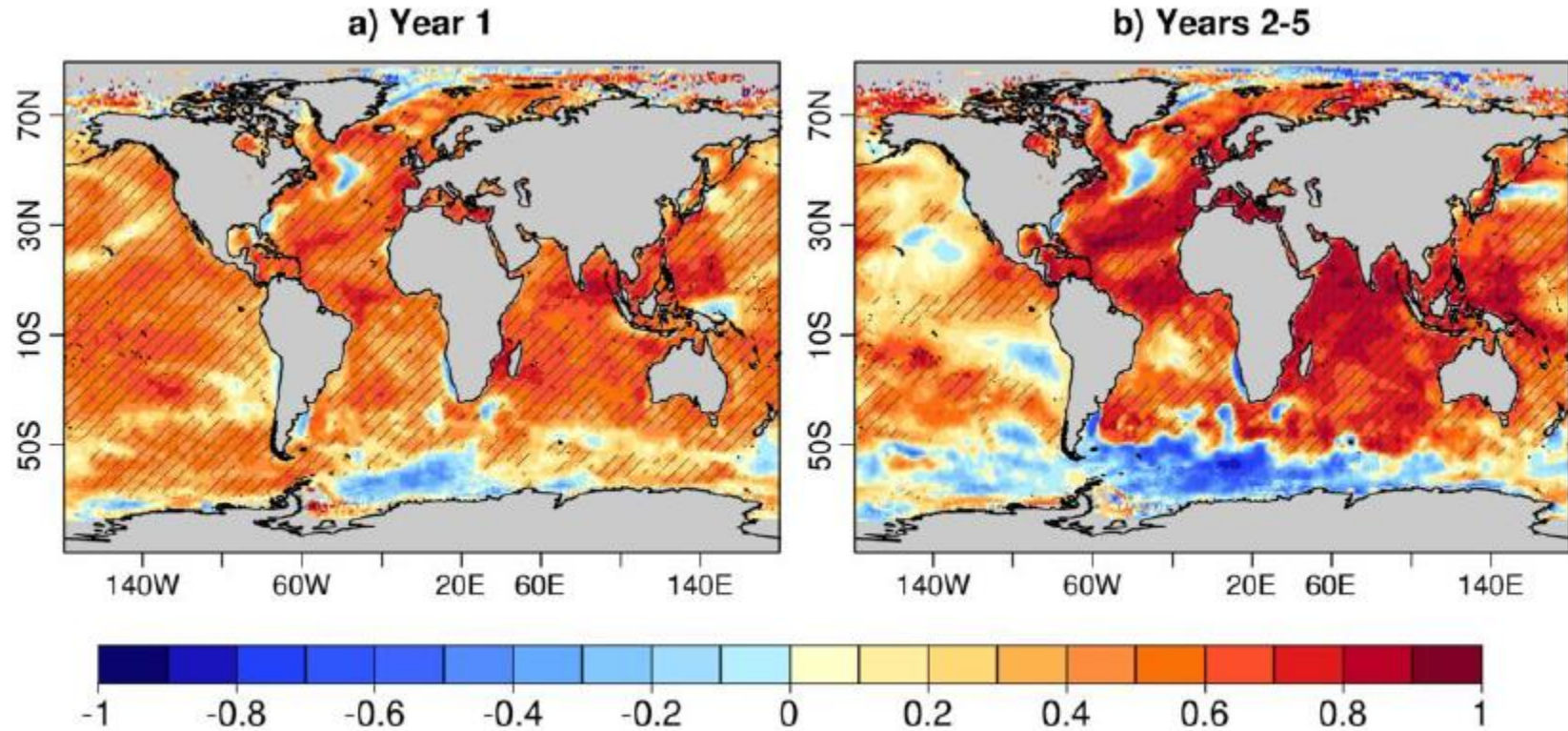


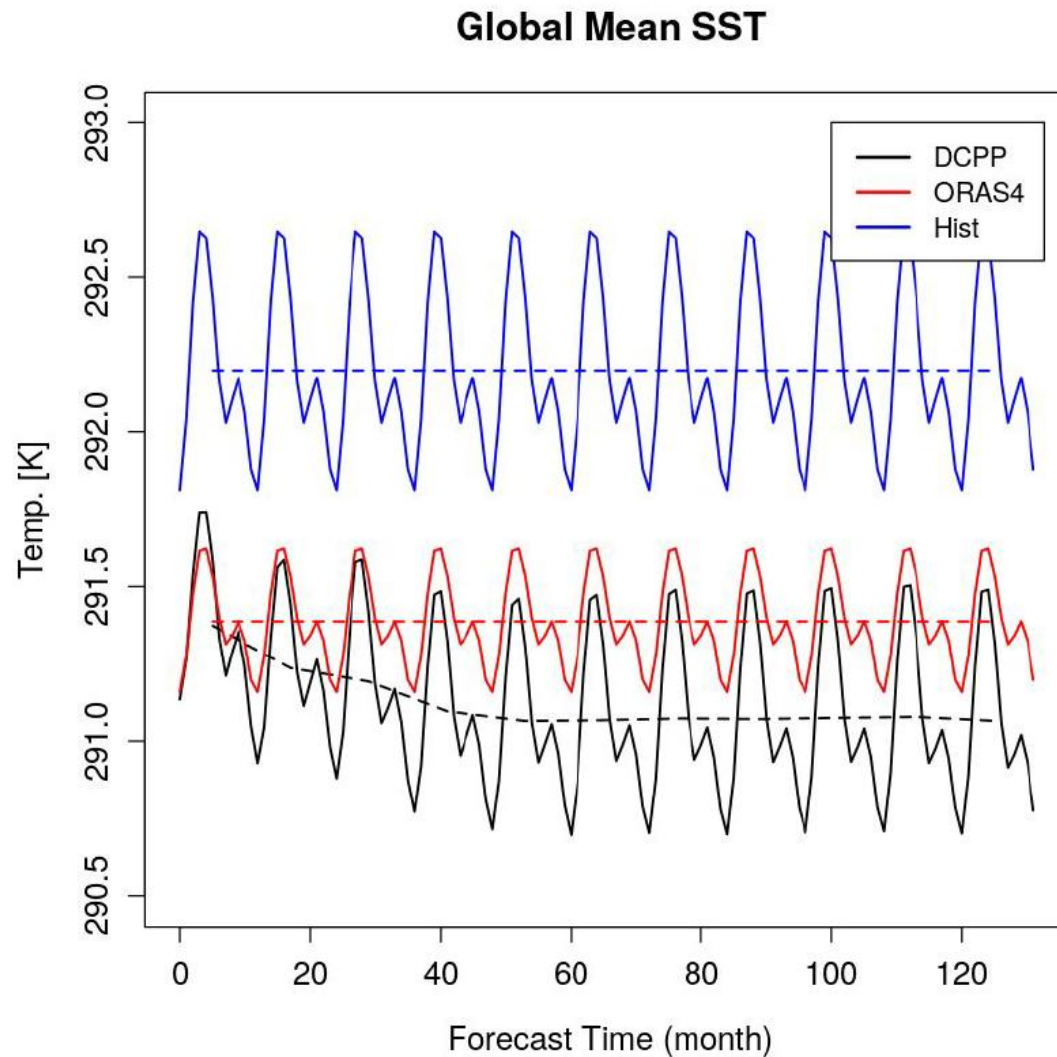
ACC Global Mean SST



After a linear detrending (which is a rather artificial approach) the skill of global mean temperature goes down substantially, especially after the 2nd forecast year

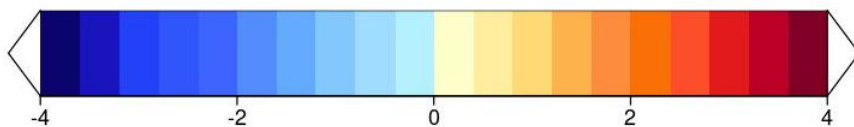
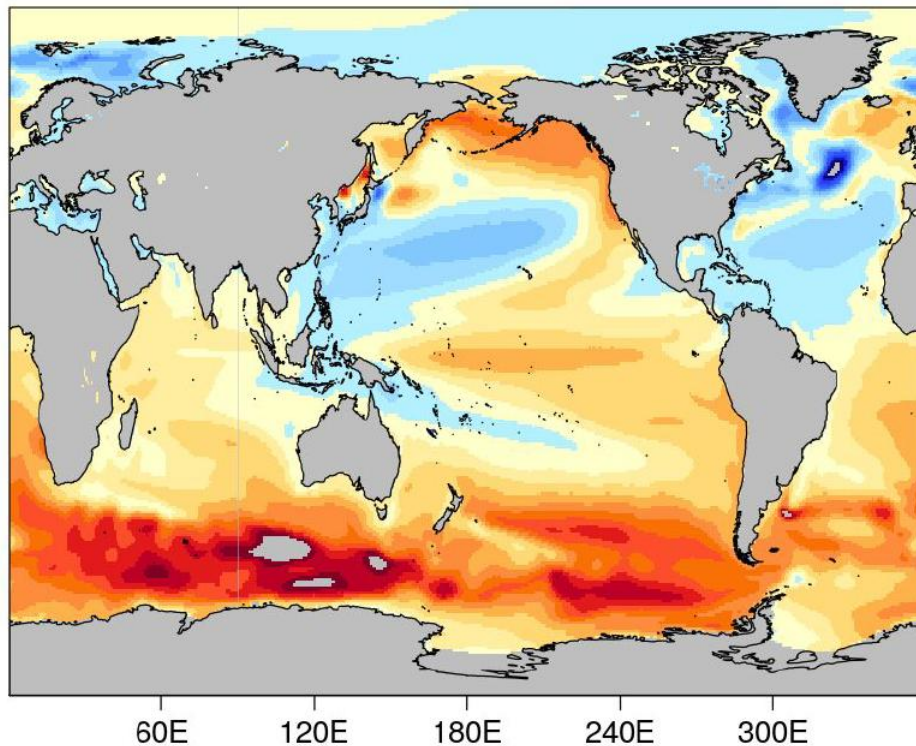
Preliminary results from DCPD Component A



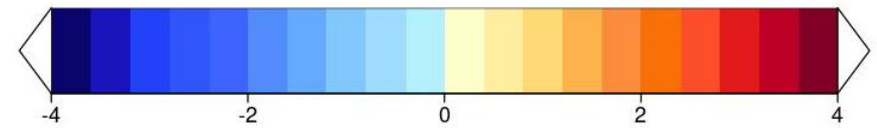
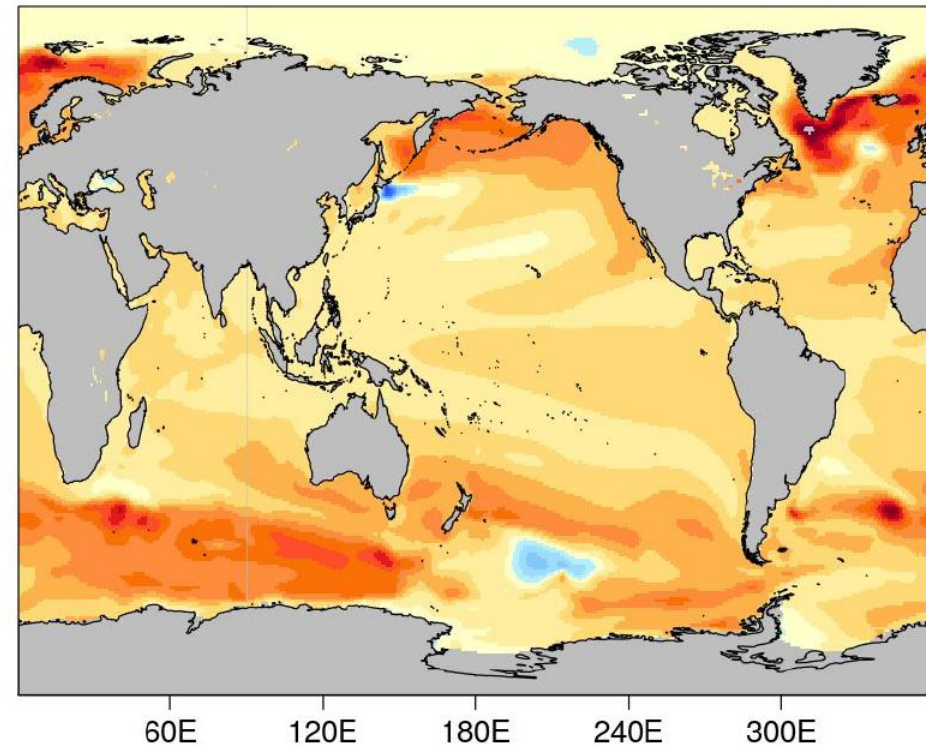


- Stronger seasonal cycle in model
- Predictions drift away from historicals
- Different attractor in historical and decadal experiments

Clim Hist - Clim DCPD (Year1), 1970-2005

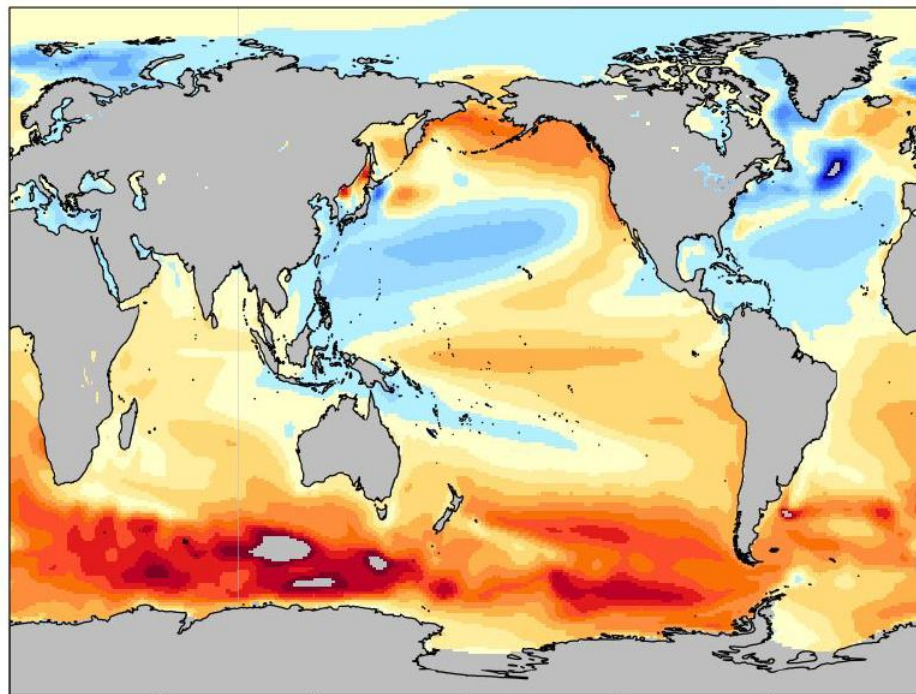


Clim Hist - Clim DCPD (Year10), 1970-2005

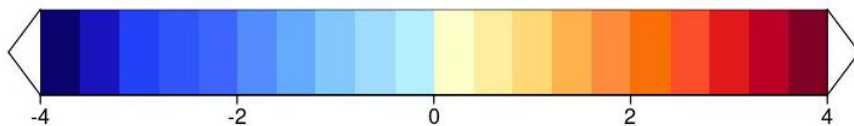


Strong warm bias in Southern Ocean, cold bias in North Atlantic

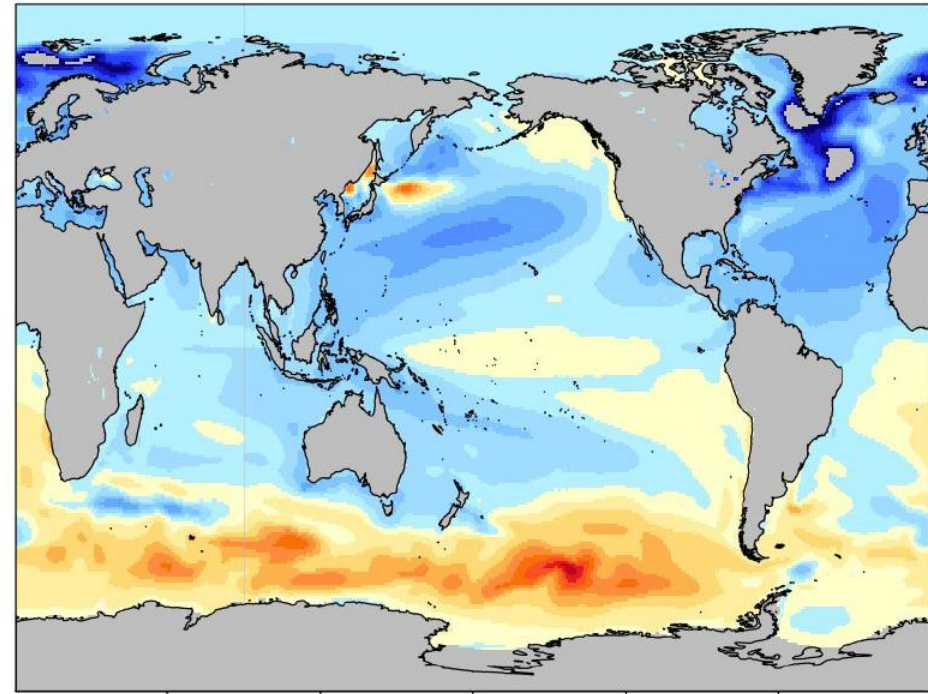
Clim Hist - Clim DCPD (Year1), 1970-2005



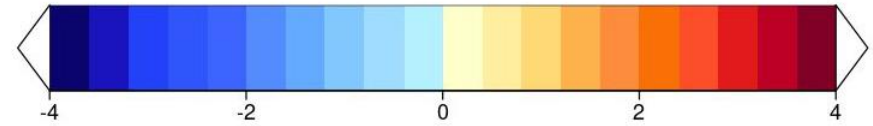
60E 120E 180E 240E 300E



Clim DCPD Yr10 - Clim DCPD Yr1, 1970-2005



60E 120E 180E 240E 300E



Weak AMOC in predictions?

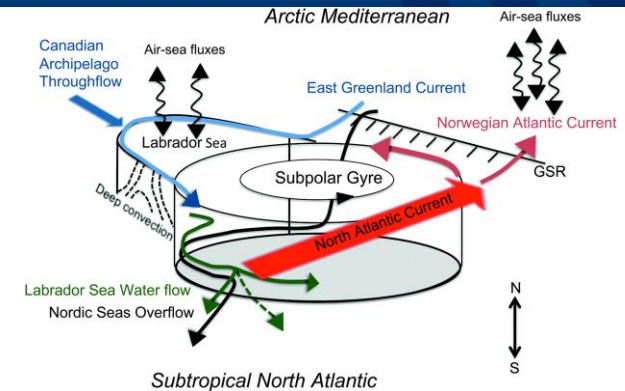
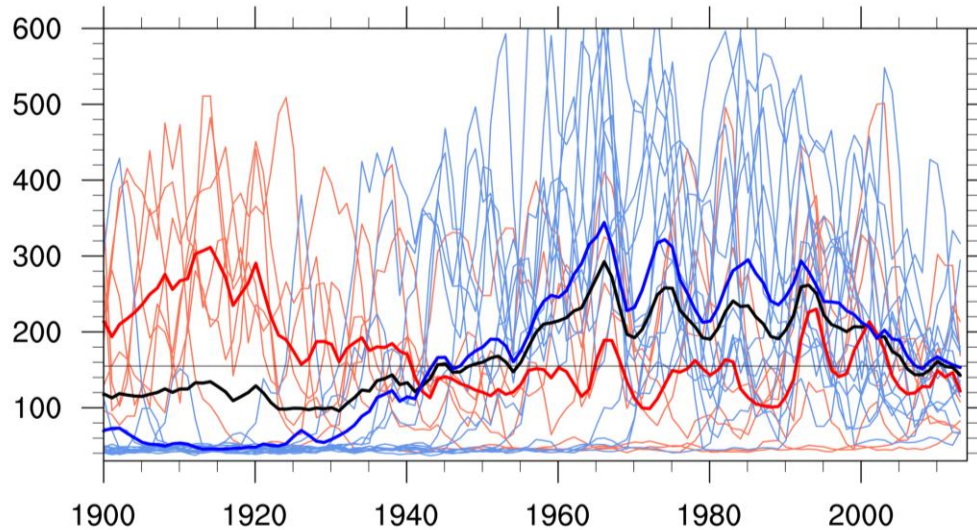
Results from Historical Ensemble



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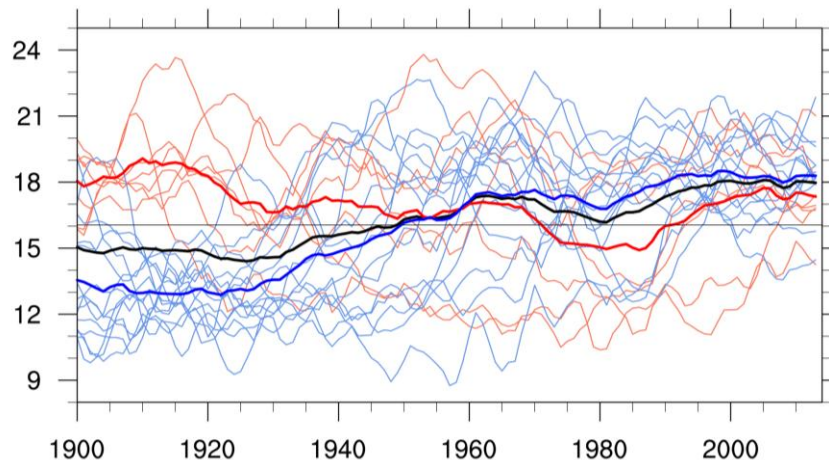
Labrador Sea mld FMA (m)



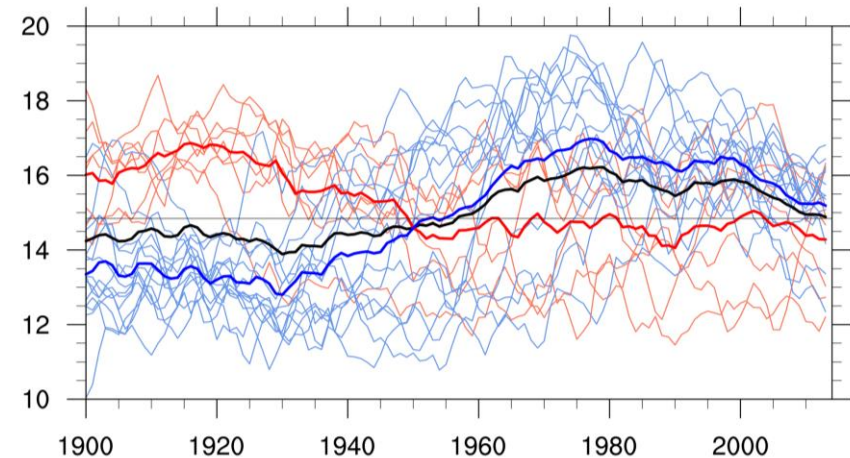
Blue members show

- no convection in Labrador Sea until 1930
- weaker Sub Polar Gyre intensity
- weaker AMOC than red members

N. Atlantic SPG Intensity JFMAMJJASOND (Sv)



AMOC@35°N-45°N/500m-1500m JFMAMJJASOND (Sv)



Results from Historical Ensemble



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**Implications for estimation of
added value of initialisation?**

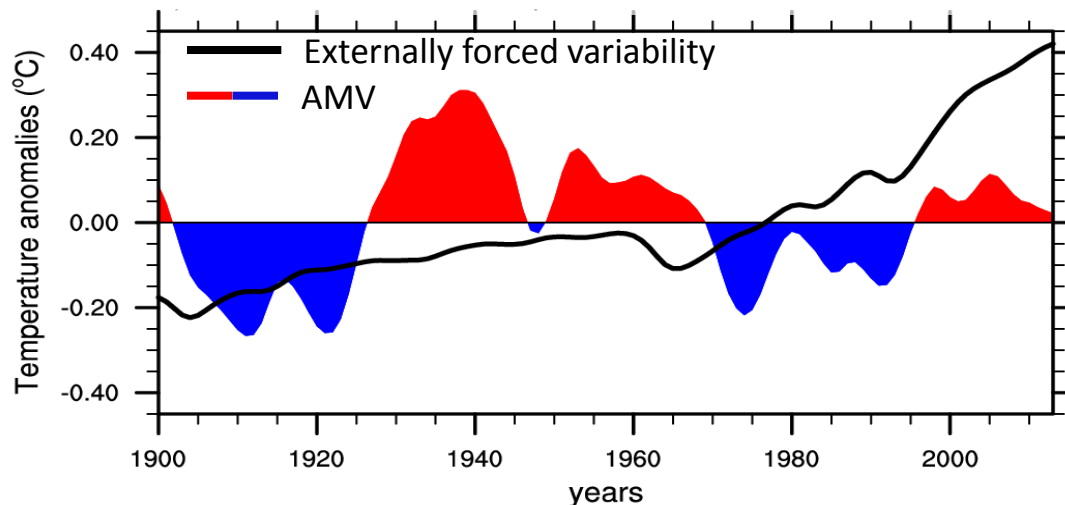
- Decadal Predictions with EC – Earth for CMIP6 DCP Component A have completed and will be on ESGF soon (DECK experiments already are)
- First (quick) results show no surprises in quality of hindcast set
→ probabilistic verification still missing
- Ongoing investigations:
 - I. What is the best way to ‘mimick’ near surface temperature observations?
 - II. Why do not initialised and initialised experiments have different model attractor?
 - III. What can we learn from behaviour of historical simulations and what implications might that have for estimation of added value of initialisation?

What am I actually doing here?

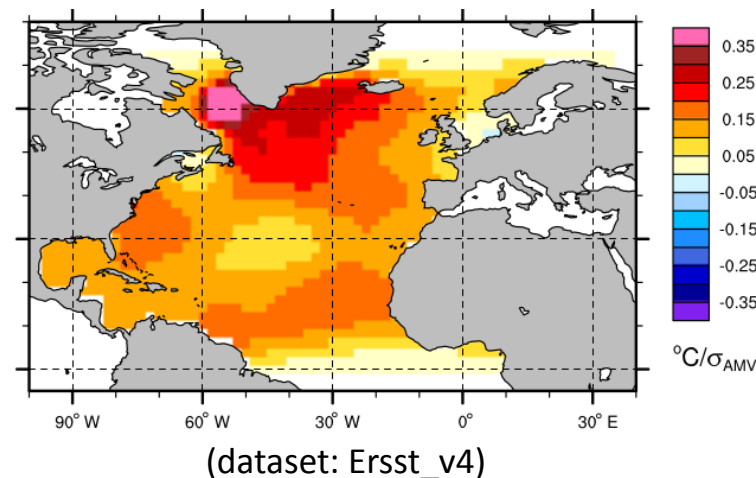
AMV impacts over Europe

- Model differences and similarities of AMV impacts on European climate
- How do pace maker experiments compare to free-running ones?

North Atlantic SST time series (Ting et al. 2009)



AMV pattern



Coupled model daily North Atlantic SST  **Climatology +/- AMV pattern**

Restoring (like a spring) of SST through non-solar surface fluxes :

$$Q_k = Q_k^0 + \gamma(SST_k^0 - SST_{AMV})$$

- Restoring coefficient of $\gamma = -40\text{W/m}^2/\text{K}$ over North Atlantic (Eq-70°N)
- Free ocean-ice-land-atmosphere interactions outside of North Atlantic

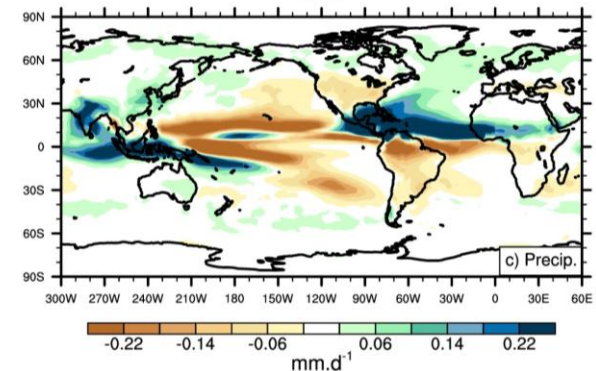
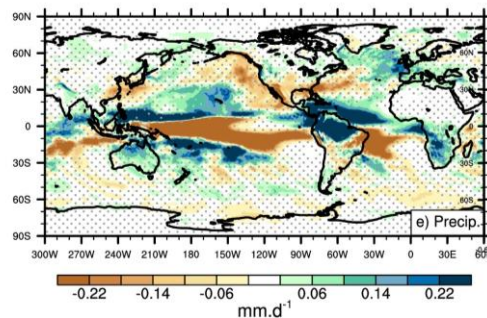
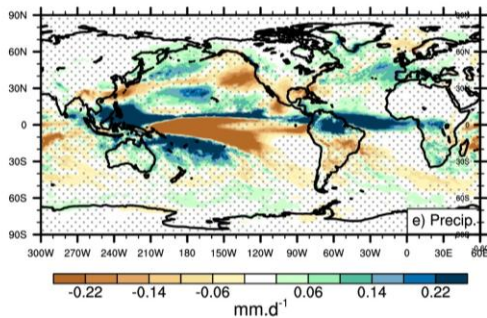
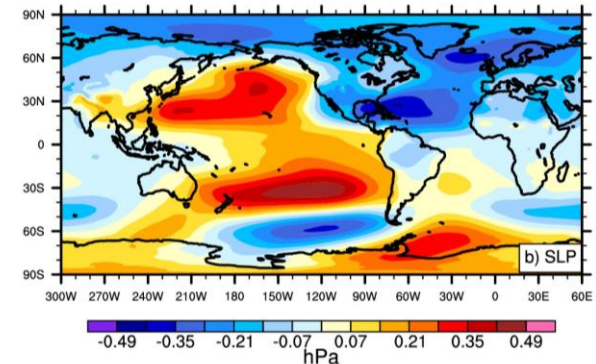
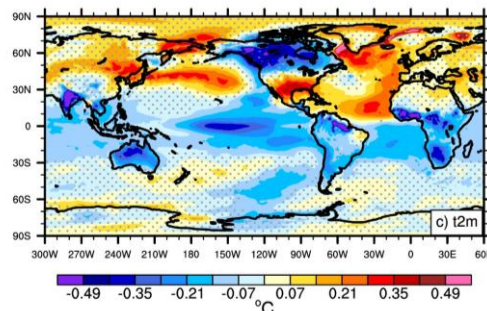
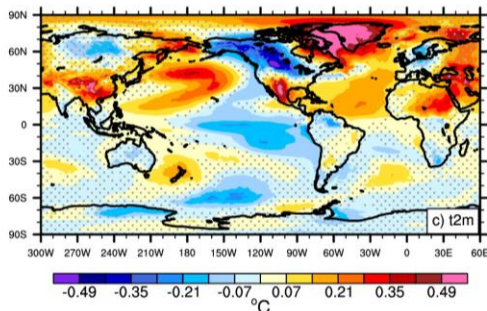
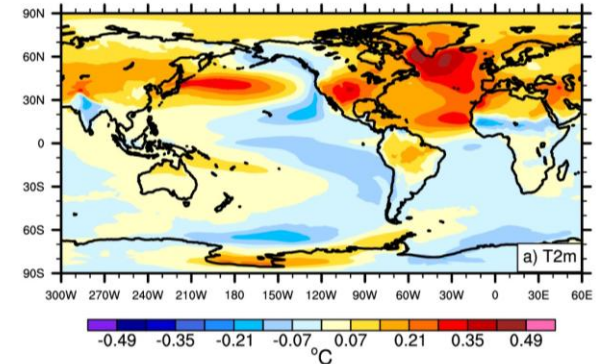
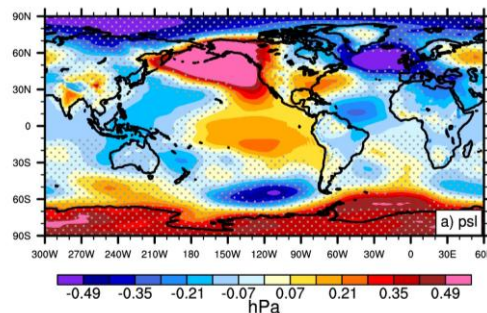
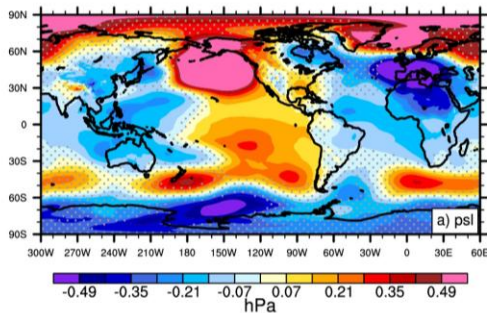
Pre DCP C Experiments



ECMWF IFS HR
AMV+ minus AMV-
DJFM

MPI ESM 1-2 HR
AMV+ minus AMV-
DJFM

MultiModel AMV+ - AMV- JJAS



Thank You!

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